

Late Permian CHIME ages of the Hida Gneiss and early Triassic age of the Mizunashi Granite in the Amo area of the Hida terrane, central Japan

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ABSTRACT

The CHIME (chemical Th-U-total Pb isochron method) dating of the Hida Gneiss, the Mizunashi Granite and a pegmatite vein have been carried out by precise microprobe analyses of monazite. The CHIME ages are 249 ± 5 and 251 ± 7 Ma for the Hida Gneiss, 237 ± 6 Ma for the Mizunashi Granite and 180 ± 3 Ma for the Pegmatite vein. There is no age signature older than ca. 250 Ma even in the center of monazite grains; monazite ages date the first attainment of Hida metamorphism to the grade of monazite formation, the amphibolite facies. The emplacement of the Mizunashi Granite at 237 Ma may correspond to the peak metamorphism of the Hida Gneiss. The pegmatite vein in the Mizunashi Granite can be correlated in age with the Jurassic Funatsu granitic rocks.

INTRODUCTION

The Hida terrane located in the northern part of central Japan consists mainly of the Hida Gneiss, the Unazuki Schist and a series of granitoids. Although there has been extensive field, petrological and geochronological studies on rocks in the Hida terrane, a controversy exists on the age of the metamorphic and granitic rocks. Up to now, the authorized facts are: (1) among the granitoids, the Funatsu Granitic Rocks are the most widespread and were emplaced in Jurassic time (ca. 180 Ma, Shibata and Nozawa, 1984), (2) the Unazuki Schist formed in late Permian or early Triassic from upper Carboniferous to Permian sediments (Hiroi, 1981), and (3) the radiometric ages of the metamorphic rocks concentrate around 240 Ma (Shibata et al., 1970; Ota and Itaya, 1989).

The main point of arguments is whether the Hida Gneiss are of true Precambrian origin or not. Minato et al. (1965) pointed out that the Hida Gneiss bears the closest resemblance to the middle Precambrian Matenrei (Macho'll-yong) System in the northeastern part of the Korean Peninsula. Sato et al. (1967) reported preliminary Rb-Sr model ages of 680-1200 Ma for the Gray

Granite that cuts the Hida Gneiss. These ages have never been confirmed, but the age data coupled with the mode of occurrence of the Gray Granite have long supported the idea that the Hida Gneiss is of polymetamorphosed Precambrian origin (e.g. Sato, 1968; Suwa, 1969,1990, Suzuki, 1977; Hiroi, 1981). Subsequent Rb-Sr studies of the Gray Granite yielded whole-rock model age as old as ca. 1100 Ma (Shibata and Nozawa, 1986) and a whole-rock isochron age of 506 Ma (Shibata et al., 1989), whereas 230-250 Ma monazite and zircon ages were reported by the CHIME method (Suzuki and Adachi, 1991b); the timing of metamorphism and plutonism still remains controversial. This paper presents the results of CHIME (chemical Th-U-total Pb isochron method) dating on the Hida Gneiss and related Mizunashi Granite from the Amo area in the central Hida terrane.

GEOLOGICAL OUTLINE

Southwest Japan is divided into Inner and Outer Zones on the basis of their position with respect to the Median Tectonic line (MTL in Fig. 1). The

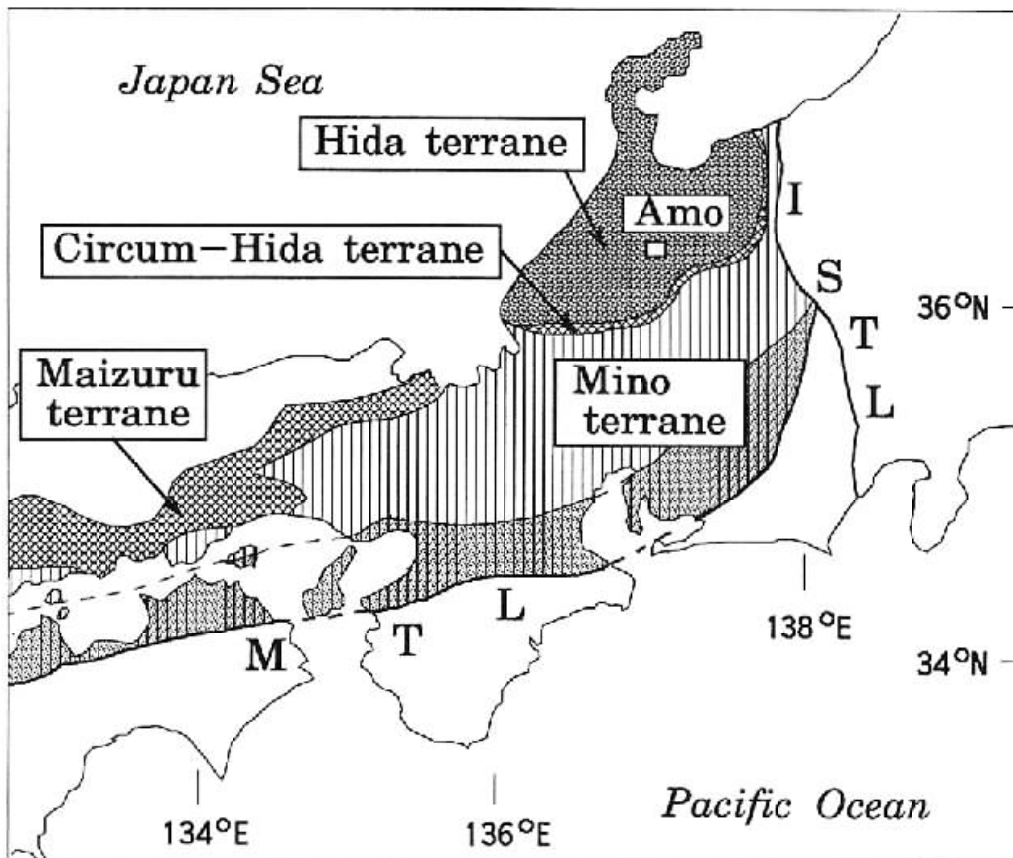


Fig. 1. Geologic framework of central Japan. MTL: the Median Tectonic Line, ISTL: the Itoigawa-Shizuoka Tectonic Line.

Inner Zone in central Japan consists of three pre-Tertiary terranes from north to south; (1) the Hida terrane of high T/P metamorphic rocks and granitoids, (2) the circum-Hida terrane of serpentinite melange with blocks of Paleozoic sediments and high P/T type metamorphic rocks, and (3) the Mino terrane of Jurassic sedimentary complex with disorganized blocks of Paleozoic limestone and greenstone and Triassic chert; the southern part of the Mino terrane is underlain by high T/P metamorphic rocks and granitoids, the Ryoike belt. The Hida terrane is underlain by the Hida Gneiss on the northwestern side, the Unazuki Schist on the southeastern side, and a variety of granitoids. They are covered with the late Jurassic to early Cretaceous Tetori Group, and Cenozoic sediments and volcanics.

Samples of the present study were collected from the Amo area. The geological configuration of the area is given in Fig. 2. The Hida Gneiss consists mainly of quartzo-feldspathic gneiss and crystalline limestone with subordinate amphibolite, pelitic gneiss and lime-silicate gneiss. The granite exposed in the western part of the study area is named as the Mizunashi Granite, and that in the southeastern part is correlatable with the Jurassic Funatsu granitic rocks (Nozawa et al., 1975).

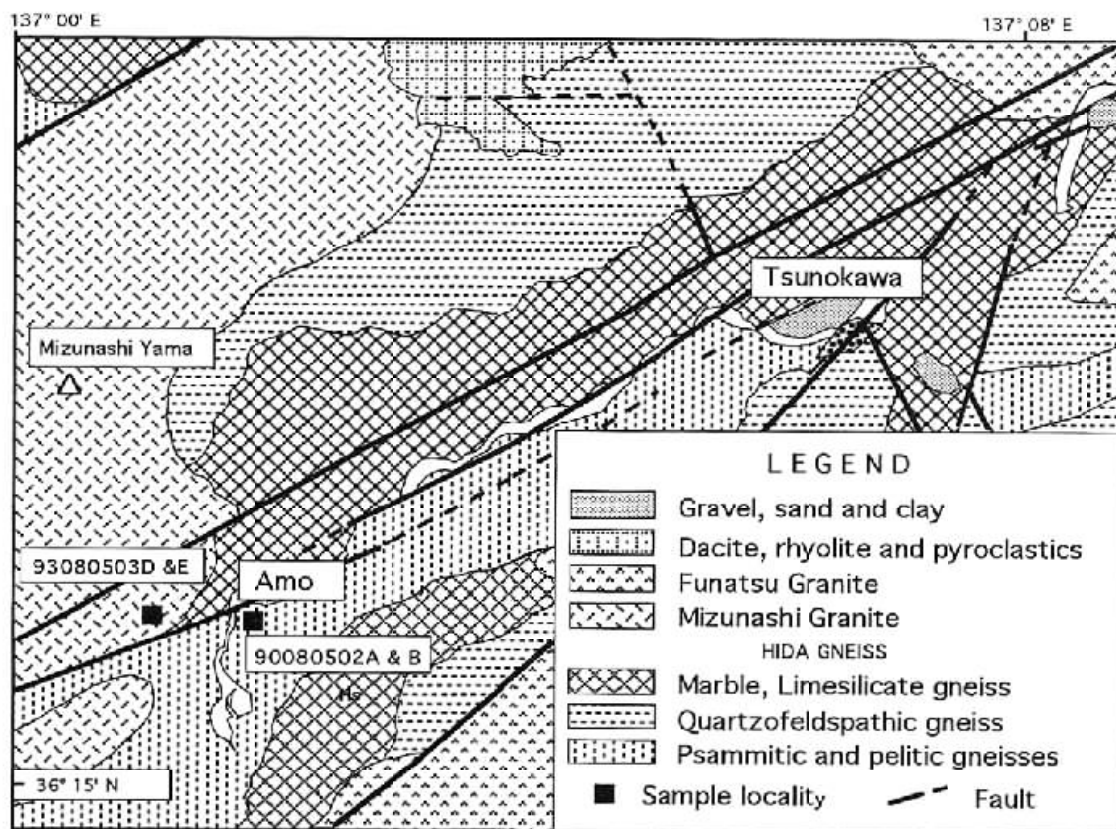


Fig. 2. Geologic map of the Amo area of the Hida terrane (simplified and modified from Nozawa et al., 1975), showing sample localities.

Two gneiss samples (Nos. 90080502A and 90080502B) were collected from thin pelitic layers intercalated with quartzo-feldspathic gneiss in an old graphite mine near Amo. The sample of Mizunashi Granite (93080503E) was collected from an outcrop about 1 km west of the mine, and sample 93080503D was collected from a pegmatite vein that cuts the Mizunashi Granite in the same outcrop as 93080503E. The gneiss samples are micaceous and consist essentially of biotite, muscovite, garnet, plagioclase and quartz with a substantial amount of graphite. Accessories include zircon, monazite, apatite and iron ores. The sample of the Mizunashi Granite is medium-grained hornblende-bearing biotite granite, and the pegmatite sample consists mainly of coarse-grained quartz and microcline with minor plagioclase, biotite and muscovite.

EXPERIMENTAL

Samples were crushed on a stamp mill. The size reduction was undertaken only to the extent necessary for disaggregation of monazite through periodic sieving to remove undersized (-60 mesh) particles before further reduction of oversized ones. Monazite and other heavy minerals were concentrated from each sieved sample by a pan. The concentrates were washed with dilute HCl and purified on an isodynamic magnetic separator. Monazite grains were hand-picked, mounted on a glass slide with Petropoxy 153 (Fig. 3A), and dried on a hot plate at 150°C. After covering with more Petropoxy 153 (Fig. 3B) and drying, the samples were polished with diamond paste until the grains were thinned approximately to half in thickness (Fig. 3C).

Monazites were analyzed on a JEOL JXA-733 electron microprobe provided with three wavelength-dispersive spectrometers (diameter of Rowland circle=280 mm). The instrument operating conditions were 15 kV accelerating voltage, 0.15 μ A probe current on Faraday cage and 5 μ m probe diameter. Intensities of ThM α , UM β and PbM α lines only were measured using PET crystal. The background was measured at two optimum positions on both sides of each line peak position. Random error due to counting statistics was deduced by counting 300s at each line peak and two background. The line and background measurements were repeated twice, and the arithmetic average of the readings was taken. The comparison standards were euxenite provided by Smellie et al. (1978) for Th and U, and synthesized glass (56.17% PbO, 13.65% ZnO and 30.18% SiO₂, analyst: K. Hayashi) for Pb.

The Th, U and Pb intensities were converted into concentrations through the Bence-Albee correction method using an average monazite composition. The detection limit of Pb (as PbO) at 2 σ confidence level is 0.007%, and the relative error in the Pb determination is around 8% at the 0.1% PbO level and much better for higher concentrations. The relative errors in the Th and U determinations are about 2% and 5%, respectively. Since the CHIME calculation has been described elsewhere (Suzuki and Adachi, 1991a,b, 1994; Adachi and Suzuki, 1992; Suzuki et al., 1991,1992,1994), the readers, if necessary, are recommended to refer to them.

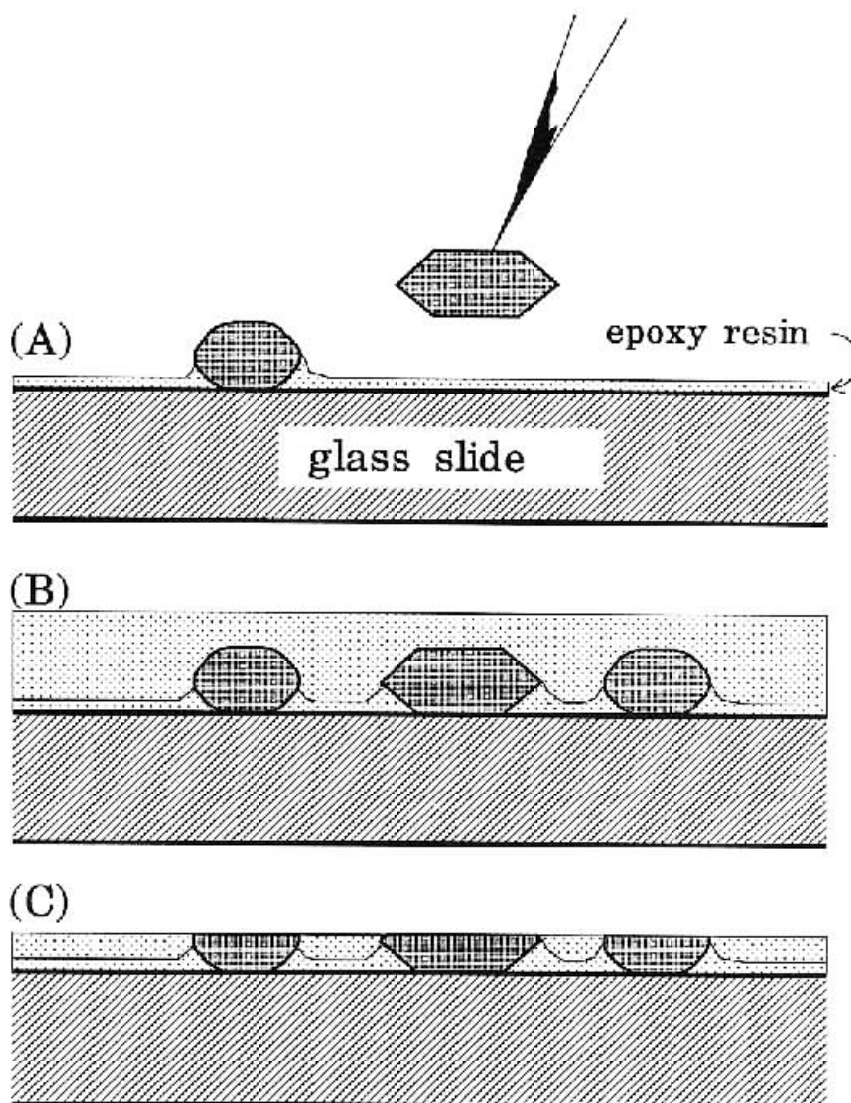


Fig. 3. Mounting of monazite grains on a glass slide (A), covering with petropoxy 154 (B), and polishing with arundum and diamond paste (C).

RESULTS

The ThO_2 , UO_2 and PbO analytical data of monazites together with apparent ages and calculated ThO_2^* concentrations are listed in Table 1.

Hida Gneiss (Samples 90080502A and 90080502B)

A total of 95 spots on 29 monazite grains from sample 90080502A were analyzed. The ThO_2^* content varies from 3.98 to 11.4%, and the PbO content from 0.044 to 0.121%. The apparent ages for most spots concentrate around 250 Ma, but those of spots 12-1 and 20-1 on metamict parts are about 200 Ma. The PbO - ThO_2^* plots of 93 data are given in Fig. 4A. The regression line, having a

Table 1. (continued).

Grain#	ThO ₂	UO ₂	PbO	age	ThO ₂ *	Grain#	ThO ₂	UO ₂	PbO	age	ThO ₂ *	Grain#	ThO ₂	UO ₂	PbO	age	ThO ₂ *
Mizunashi Granite (Sample 93080503E)						M06-02	8.82	0.061	0.092	240	9.02	M06-02	1.58	0.022	0.014	195	1.65
M01-01	10.7	0.216	0.117	242	11.4	M06-03	7.62	0.022	0.078	239	7.69	M06-03	1.62	0.079	0.014	177	1.87
M01-02	6.72	0.121	0.072	238	7.11	M06-04	7.30	0.030	0.074	235	7.40	M07-01	2.62	0.014	0.019	164	2.66
M01-03	5.90	0.074	0.064	246	6.14	M06-05	11.5	0.026	0.116	236	11.6	M07-02	2.69	0.029	0.021	176	2.78
M01-04	3.16	0.069	0.035	247	3.38	M06-06	4.88	0.070	0.051	236	5.11	M08-01	2.79	0.021	0.021	171	2.86
M01-05	8.01	0.100	0.086	244	8.33							M08-02	1.70	0.020	0.012	166	1.76
M01-06	7.55	0.071	0.078	238	7.78							M08-03	1.67	0.023	0.013	176	1.74
M01-07	8.96	0.055	0.096	247	9.14	Pegmatite (sample 93080503D)						M08-04	2.31	0.037	0.018	170	2.43
M02-01	8.62	0.128	0.092	241	9.03	M01-01	2.45	0.023	0.019	181	2.52	M09-01	2.61	0.021	0.021	185	2.68
M02-02	8.79	0.138	0.098	251	9.24	M01-02	2.07	0.016	0.018	196	2.12	M10-01	2.29	0.014	0.018	183	2.33
M02-03	8.78	0.133	0.093	239	9.21	M01-03	1.84	0.013	0.015	188	1.88	M10-02	3.08	0.025	0.024	180	3.16
M02-04	13.3	0.153	0.139	239	13.8	M02-01	2.49	0.035	0.019	176	2.60	M10-03	1.18	0.165	0.012	168	1.71
M02-05	9.27	0.139	0.097	236	9.72	M02-02	2.05	0.019	0.015	173	2.11	M10-04	4.38	0.138	0.038	188	4.82
M02-06	13.7	0.178	0.143	236	14.3	M02-03	1.84	0.017	0.015	181	1.90	M10-05	2.46	0.023	0.020	186	2.53
M02-07	12.2	0.181	0.132	245	12.8	M02-04	2.41	0.027	0.019	183	2.50	M10-06	4.44	0.024	0.034	178	4.52
M02-08	10.7	0.181	0.115	241	11.3	M02-05	2.46	0.018	0.020	189	2.52	M10-07	2.35	0.037	0.018	171	2.47
M03-01	10.4	0.304	0.112	233	11.4	M02-06	1.90	0.015	0.014	165	1.95	M10-08	4.95	0.108	0.041	183	5.30
M03-02	11.2	0.291	0.120	233	12.1	M03-01	2.84	0.030	0.022	178	2.94	M11-01	5.94	0.116	0.048	179	6.32
M03-03	12.9	0.295	0.143	243	13.9	M03-02	2.02	0.039	0.015	162	2.15	M11-02	3.10	0.030	0.026	188	3.20
M03-04	10.9	0.285	0.118	236	11.8	M03-03	1.95	0.034	0.016	188	2.06	M11-03	4.43	0.042	0.035	183	4.57
M03-05	11.0	0.305	0.124	245	12.0	M03-04	1.71	0.017	0.013	174	1.76	M11-04	2.99	0.025	0.024	182	3.07
M03-06	10.1	0.304	0.114	243	11.1	M03-05	1.86	0.020	0.014	168	1.92	M11-05	4.50	0.035	0.037	191	4.61
M04-01	10.1	0.093	0.105	237	10.4	M04-01	13.9	0.140	0.109	180	14.4	M11-06	4.09	0.038	0.030	169	4.21
M04-02	7.23	0.108	0.077	241	7.58	M04-02	15.1	0.185	0.115	174	15.7	M12-01	2.30	0.015	0.017	175	2.35
M04-03	10.2	0.122	0.106	235	10.6	M04-03	15.0	0.199	0.122	184	15.6	M12-02	2.07	0.037	0.016	171	2.19
M04-05	10.1	0.128	0.105	236	10.5	M04-04	17.5	0.160	0.136	178	18.0	M12-03	4.34	0.053	0.034	180	4.51
M05-01	9.76	0.106	0.100	235	10.1	M04-05	12.0	0.059	0.092	178	12.2	M12-04	5.29	0.083	0.043	182	5.56
M05-02	9.35	0.103	0.097	236	9.68	M04-06	31.1	0.294	0.244	180	32.0	M12-05	4.40	0.340	0.039	166	5.50
M05-03	12.1	0.140	0.127	239	12.6	M04-07	15.0	0.130	0.117	179	15.4	M12-06	7.49	0.059	0.059	180	7.68
M05-04	12.9	0.081	0.129	233	13.2	M05-01	4.56	0.026	0.034	174	4.64	M13-01	2.30	0.024	0.018	175	2.38
M05-05	12.5	0.078	0.129	238	12.8	M05-02	3.63	0.024	0.029	186	3.71	M13-02	3.20	0.013	0.025	181	3.24
M05-06	8.95	0.070	0.092	237	9.18	M05-03	3.70	0.051	0.028	172	3.86	M13-03	3.44	0.019	0.026	178	3.50
M06-01	6.35	0.023	0.064	234	6.42	M06-01	3.74	0.018	0.030	187	3.80	M13-04	4.74	0.030	0.036	177	4.84

slope of 0.0105 ± 0.0003 and an intercept value of 0.0007 ± 0.0023 , yields an age of 250.5 ± 6.8 Ma (MSWD=0.17).

A total of 248 spots on 47 monazite grains from sample 90080502B were analyzed. The ThO₂* content varies from 5.47 to 14.4% and the PbO content from 0.057 to 0.149%. The apparent ages are in the range between 240 and 261 Ma, but three spots (M01-04, M04-33 and, M04-34) on metamict parts give much younger apparent ages of 200-180 Ma. The PbO-ThO₂* plots of 245 data are given in Fig. 4B. They are regressed with an isochron of 248.9 ± 4.5 Ma (MSWD=0.13) with a slope of 0.0105 ± 0.0002 and an intercept value of 0.0005 ± 0.0019 .

Mizunashi Granite (Sample 93080503E)

A total of 37 spots on transparent portions of 6 monazite grains were analyzed. Monazite contains 3.38 to 14.3% ThO₂* and 0.0353 to 0.1428% PbO. All data points are arrayed linearly on the PbO vs. ThO₂* diagram (Fig. 5A), yielding isochron (MSWD=0.12) of 237 ± 5 Ma with a slope of 0.0100 ± 0.0002 and an intercept value of 0.0009 ± 0.0020 .

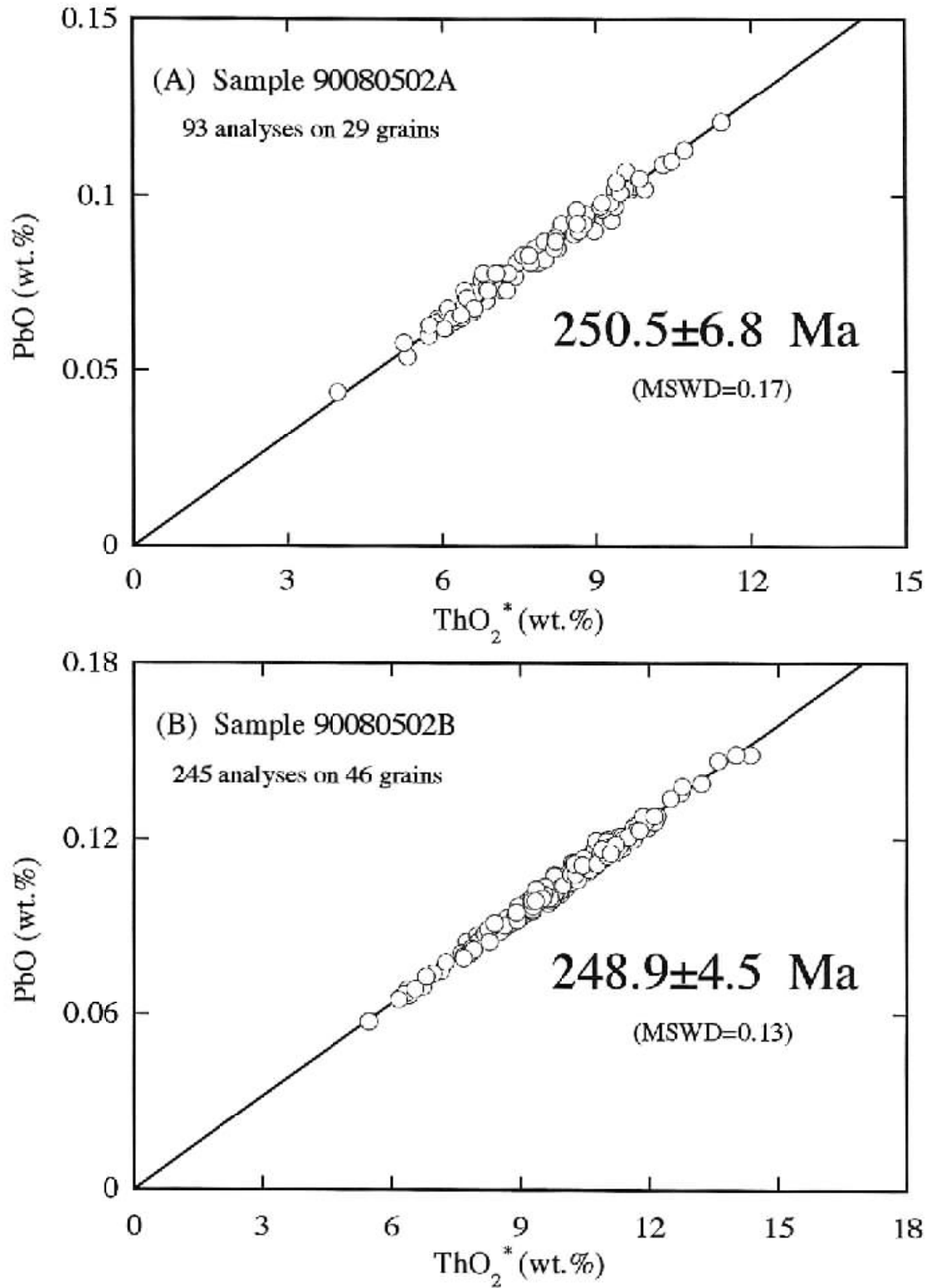


Fig. 4. Plots of PbO vs. ThO_2^* of monazite grains from sample 90080502A (A) and from sample 90080502B (B). The error given to the age calculation is of 2σ .

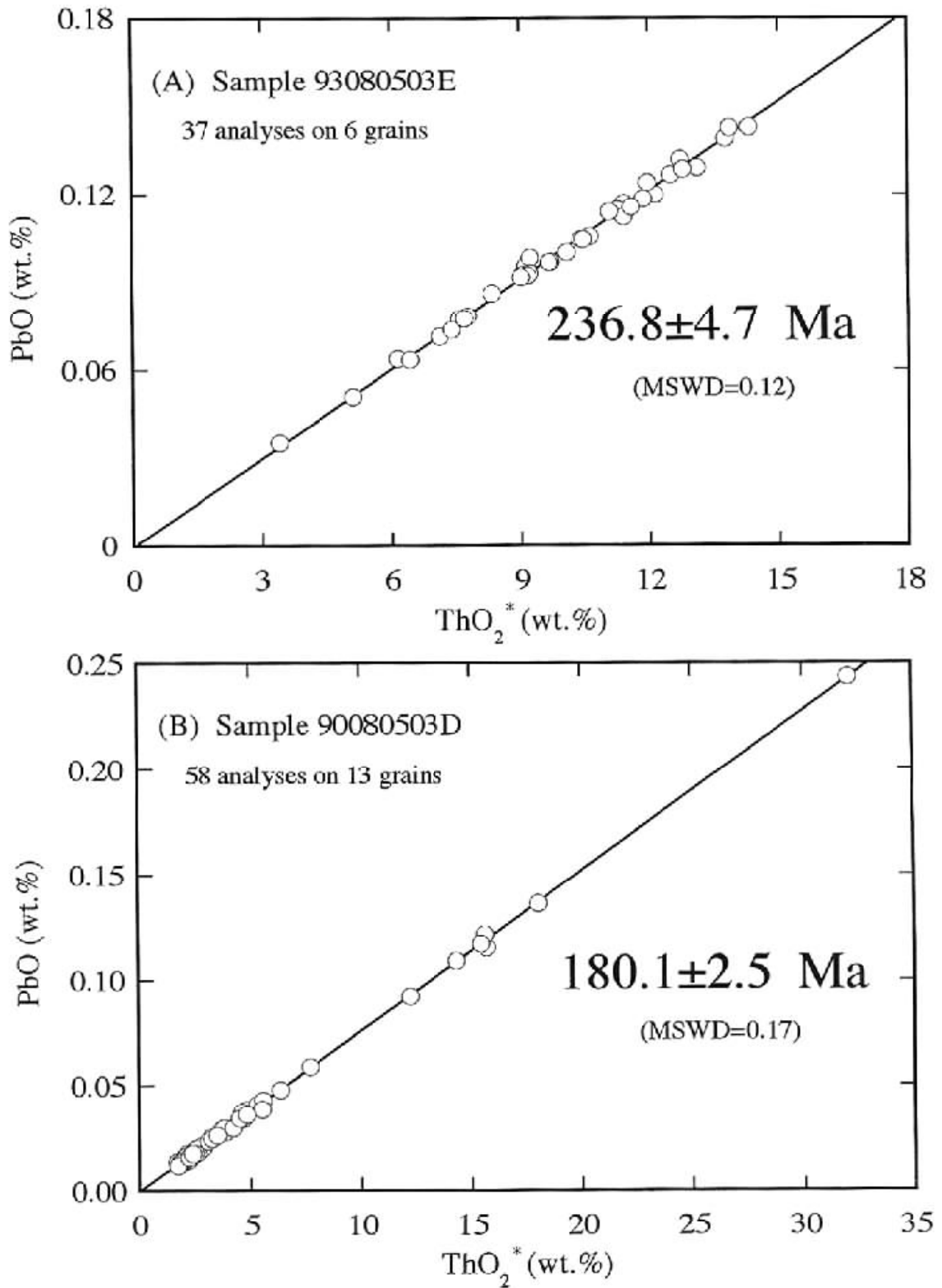


Fig. 5. Plots of PbO vs. ThO₂* of monazite grains from the Mizunashi Granite (A) and a pegmatite vein (B). The error given to the age calculation is of 2 σ .

Pegmatite vein (Sample 93080503D)

A total of 58 spots on transparent portions of 13 monazite grains were analysed. The ThO₂* content varies from 1.71 to 32.0%, and the PbO content from 0.012 to 0.244%. The apparent ages for individual spots are in the range between 169 and 196 Ma. The PbO-ThO₂* plots of 58 data are given in Fig. 5B. The regression line, having a slope of 0.0076 ± 0.0001 and an intercept value of -0.0001 ± 0.0004 , yields an age of 180.1 ± 2.5 Ma (MSWD=0.17).

DISCUSSION

The CHIME monazite ages for samples of the Hida Gneiss, 250.5 ± 6.8 and 248.9 ± 4.5 Ma, agree well with each other. The closure temperature of Pb diffusion in monazite is 650-700 °C (Suzuki et al., 1994). Smith and Barreiro (1970) found that monazite in gneisses formed as metamorphic mineral at the lower amphibolite facies conditions records the time since the formation even with subsequent excursions into upper amphibolite facies. The stable association of muscovite and quartz shows that the gneiss samples did not experience the upper amphibolite facies conditions. We thus consider that the ca. 250 Ma monazite ages date the first attainment to the lower amphibolite facies grade during the prograde stage of the Hida metamorphism.

If the Hida Gneiss is of polymetamorphosed Precambrian origin as has been speculated previously, monazite grains should bear signature of older ages. Despite our search for Precambrian monazite cores over 29 grains from sample 90080502A and 46 grains from sample 90080502B, no parts older than ca. 250 Ma have been found. Suzuki and Adachi (1994) reported the occurrence of ca. 350 Ma detrital zircons from psammitic gneiss samples in the Tsunokawa area, east of the Amo area. This indicates that sedimentation of the gneiss protolith took place in the middle-late Paleozoic. Again, we can state that the Hida metamorphic rocks in the Amo area formed in the late Permian, and are not of polymetamorphosed Precambrian rocks.

The CHIME monazite age of the Mizunashi Granite is 236.8 ± 4.7 Ma. This age is slightly younger than the previously reported Rb-Sr whole-rock isochron age of 296.7 ± 25.6 Ma (Shibata and Nozawa, 1984). Since the Rb-Sr isotopic data are scattered, we here regard the 236.8 ± 4.7 Ma age as the time of granite emplacement. Another Rb-Sr mineral isochron age for the Mizunashi Granite was reported as 210.9 ± 1.8 Ma (Shibata and Nozawa, 1984). This age is evidently older than the ages for the post-tectonic Funatsu granitic rocks (188.9 ± 4.4 Ma; Shibata and Nozawa, 1984), and is correlatable with K-Ar hornblende and muscovite ages for the Hida Gneiss. Therefore, it seems much more likely that the Mizunashi Granite was emplaced possibly at the peak metamorphism and cooled together with the Hida Gneiss to the blocking temperatures of hornblende and muscovite at ca. 210 Ma.

The CHIME monazite age of 180.1 ± 2.5 Ma for the pegmatite vein cutting the Mizunashi Granite is correlative with the 188.9 ± 4.4 Ma Rb-Sr whole-rock isochron age of the Funatsu granitic rocks (Shibata and Nozawa, 1984). Probably

the pegmatite vein had been derived from the late stage residual of the Funatsu granitic rocks.

CONCLUDING REMARKS

The CHIME ages of monazite have revealed that the Hida Gneiss in the Amo area originated from late Paleozoic sediments and attained at the lower amphibolite facies grade in late Permian time (ca. 250 Ma). The peak metamorphism was possibly contemporaneous with the emplacement of the Mizunashi Granite at ca. 237 Ma. Taking available K-Ar and Rb-Sr mineral ages into consideration, we can conclude that both the gneisses and granite cooled to blocking temperature of hornblende and muscovite at ca. 210 Ma. The Hida Gneiss and the Mizunashi Granite subsequently underwent the intrusion of pegmatite at ca. 180 Ma. The previous view that Hida metamorphic rocks formed in Precambrian time is no longer valid.

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